Donald Rapp

Ice Ages and Interglacials

Measurements, Interpretation, and Models

Third Edition



Ice Ages and Interglacials

Donald Rapp

Ice Ages and Interglacials

Measurements, Interpretation, and Models

Third Edition



Donald Rapp South Pasadena, CA, USA

ISBN 978-3-030-10465-8 ISBN 978-3-030-10466-5 (eBook) https://doi.org/10.1007/978-3-030-10466-5

Library of Congress Control Number: 2018965882

1st edition: © Praxis Publishing Ltd, Chichester, UK 2009

2nd edition: © Springer-Verlag Berlin Heidelberg 2012

3rd edition: © Springer Nature Switzerland AG 2019

This work is subject to copyright. All rights are reserved by the Publisher, whether the whole or part of the material is concerned, specifically the rights of translation, reprinting, reuse of illustrations, recitation, broadcasting, reproduction on microfilms or in any other physical way, and transmission or information storage and retrieval, electronic adaptation, computer software, or by similar or dissimilar methodology now known or hereafter developed.

The use of general descriptive names, registered names, trademarks, service marks, etc. in this publication does not imply, even in the absence of a specific statement, that such names are exempt from the relevant protective laws and regulations and therefore free for general use.

The publisher, the authors and the editors are safe to assume that the advice and information in this book are believed to be true and accurate at the date of publication. Neither the publisher nor the authors or the editors give a warranty, express or implied, with respect to the material contained herein or for any errors or omissions that may have been made. The publisher remains neutral with regard to jurisdictional claims in published maps and institutional affiliations.

This Springer imprint is published by the registered company Springer Nature Switzerland AG The registered company address is: Gewerbestrasse 11, 6330 Cham, Switzerland

Preface

The typical description of the past 800,000 years would be that the Earth has experienced about nine major periods of glaciation ("Ice Ages") spaced at various intervals ranging from 0.9 my to 1.1 my (see Fig. 8.23). This presupposes that Ice Ages are unusual departures from normalcy. Actually, it appears as if the natural state of the Earth during this period was an Ice Age, but there were about nine interruptions of the glacial state, during which the Arctic climate was much warmer for time periods of the order of 10,000 years or so. Each Ice Age required many tens of thousands of years to develop to its maximum state of glaciation.

During the last glacial maximum, some 20,000 years ago, Canada and the northern USA were blanketed by huge ice sheets, up to 4 km in thickness. In addition, there was a large ice sheet covering Scandinavia that reached down into Northern Europe. The Antarctic ice sheet was somewhat more extensive than today. Local glaciations existed in mountainous regions of North America, Europe, South America and Africa, driving the tree line down by up to 700–800 m. The temperature of Greenland was lower by up to 20 °C, but the climate was probably only a few degrees colder than normal in the tropics. Conditions were very harsh 20,000 years ago at the last glacial maximum (LGM).

These ice sheets tied up so much of the Earth's water that more than 150 m of ocean was removed. As a result, the shorelines of the continents moved outward by a considerable distance. The Beringia land bridge from Siberia to Alaska was created, allowing animals and humans to cross from one continent to the other. In the upper-mid latitudes, the climates were semi-arctic and the flora shifted to tundra. Humidity was reduced, and many lands dried out. At the LGM, the CO_2 concentration dropped below 200 ppm, and this combined with cold, led to plant starvation and desertification of marginal areas. The sharp temperature discontinuity at the edges of the ice sheets generated violent winds that swept up dust and dirt from dry regions, filling the atmosphere with dust. Dust deposited on the ice sheets preceded each termination of an Ice Age, allowing much greater solar absorption and the demise of the ice sheets. This last Ice Age began to wane around 17,000 years ago and dissipated through a series of gyrating climate oscillations, ending in a comparatively benign period that has lasted for the past ~10,000 years, called the *Holocene*.

A few geologists of the nineteenth century were perceptive enough to read the signs in the rocks and geological formations, and concluded that the Earth must have once (or more) been heavily glaciated with massive ice sheets that generated the markings and rock depositions that they observed. They eventually overcame the initial resistance to this new (and shocking) concept in the geological community. But it was not until the 1970s that extensive studies of marine sediments (followed by polar ice core studies in the 1980s and 1990s) demonstrated the existence, amplitude and recurrent chronology of multiple Ice Ages.

During the nineteenth century, several scientists proposed that Ice Ages could have resulted from semi-periodic variability in the Earth's orbital parameters, which change the relative solar energy input to higher latitudes. As the theory goes, when summer solar energy input to higher northern latitudes drops below a critical threshold range, ice and snow can better survive the summer. Data acquired in the twentieth century suggests that ice sheets slowly begin to form over many millennia at latitudes roughly in the range 60° N to 70°N. As the ice cover spreads, the albedo (reflectivity) of the region increases, further adding to the cooling effect. More and more water leaves the oceans and gets deposited into the building ice sheets, lowering the oceans and extending shorelines outward. Since land has a higher albedo than oceans, this provides further cooling. In the regions adjacent to the ice sheets, vegetation is inhibited, adding still further to increased Earth albedo. As the northerly regions cool, the concentrations of key greenhouse gases, water vapor, CO_2 and CH₄, decrease, adding to a worldwide cooling effect that makes the budding Ice Age a global phenomenon. Other effects such as widespread dust storms and expansion of sea ice and mountain glaciers also contributed. Thus, a runaway expansion of ice sheets develops over many millennia. James Croll formulated the concept of the solar trigger for Ice Ages based on variations of the Earth's orbit in 1875. In the first several decades of the twentieth century, M. Milankovitch quantified this theory by carrying out extensive calculations by hand in the pre-computer age. Nevertheless, in the absence of long-term data over many Ice Ages, the astronomical theory remained an abstract concept. Furthermore, there were no credible mechanistic models that described how changing solar energy inputs to higher latitudes led specifically to alternating Ice Ages and deglaciations.

With the advent of marine sediment data in the 1970s, it became possible to compare the astronomical theory with data over many glacial cycles. John Imbrie was a pioneer in this regard. He created the SPECMAP "stack" of ocean sediment data from several sites to reduce noise and devised models with which to compare ice sheet volume (v) to solar variations. Lacking an absolute dating methodology for the sediment data, he "tuned" the chronology of the SPECMAP to the variations in solar input to high altitudes. He also used spectral analysis to show that some of the prominent frequency components of the SPECMAP variability were in consonance with known frequencies of solar variation. From this, he concluded that the astronomical model explained much of the Ice Age record—at least for the past ~650,000 years. However, there is circular reasoning involved, and one could construe his procedure to involve curve fitting as well as physics. It seems clear that the solar input to high latitudes is involved in setting the timing of transitions between periods of glaciation, but there is no quantitative theory that predicts a priori when these transitions take place.

As ocean sediment data was extended backward in time, it became apparent that the glacial cycles were evidently controlled by the Sun, but the details were difficult to work out. Of greatest importance was the fact that the period from about 2.7 mya to about 800 kya was characterized by relatively rapid, smaller amplitude climate cycles, whereas since about 800 kya, climate cycles have consistently increased in period and amplitude. By contrast, the astronomical theory would not have predicted any such major shift in frequency and amplitude since there is no reason to believe that solar forcing to higher latitudes changed qualitatively during this time period. However, Raymo et al. (2006) proposed an explanation for this that makes good sense. There were other problems with astronomical theory as well; at some prominent occurrences of climate change, there were no corresponding variations in solar input (e.g., 400 kya). Since the 1990s, a number of studies have attempted to resolve the differences between the data and the astronomical theory. Some of these studies had an obvious and pervasive bias in favor of the astronomical theory against all odds.

Yet despite the problems with the astronomical theory, there are several tantalizing similarities between the climate data and the historical solar record. These include the correlation of several important frequencies in spectral analyses and certain undeniable rough similarities in the timing of climate and solar records over some periods during the past several hundred thousand years.

The main problem with astronomical theory is that it is not at all clear just what the theory is! What seems to be most glaringly absent from the astronomical theory is a clear quantitative mechanism by which variations in solar input to higher latitudes produce changes in climate. The theory seems to revolve about the notion that when solar input to high northern latitudes is high, the climate tends to be interglacial, and when solar input to high northern latitudes is low, the climate tends to be glacial. The cycles prior to about 800 kya follow a period of about 41 ky. Subsequent to about 800 kya, as the Earth grew colder, the ice sheets thickened considerably and the spacing of glacial cycles more than doubled. The evidence suggests that in this later regime, the natural state of the Earth was what we call an Ice Age. Lacking any other perturbation, the energy balance of the Earth (prior to modern industrial times) favored growth of ice sheets at high northern latitudes. Starting at any arbitrary time, the ice sheets grew for several tens of thousands of years. During this period of growth of the ice sheets, the peak midsummer solar input to high latitudes oscillated with its $\sim 22,000$ -year period due to precession. Up-lobes in solar slowed down expansion of the ice sheets, and down-lobes increased the rate of expansion of the ice sheets, but the Ice Age would persist through several of these 22,000-year precessional cycles. After perhaps four precessional cycles, when the ice sheets became very extensive and the global CO₂ concentration dropped below 200 ppm, a precessional solar up-lobe led to a rapid termination of the ice sheets. They disintegrated in a mere 5500 years or so. This led to an interglacial, which was eventually followed by gradual evolution of yet another Ice Age. Evidently, solar input to high northern latitudes is involved, because all terminations occur at up-lobes in the solar oscillation. Yet, many up-lobes during an evolving Ice Age do not produce terminations. Only for a very mature Ice Age, after perhaps four or five precessional cycles, does a solar up-lobe lead to a termination. Therefore, there must be some X-factor that is necessary to induce a rapid termination, in consonance with the up-lobe in solar input. The X-factor only occurs in a very mature Ice Age.

Ellis and Palmer (2016) noted that dust levels in ice cores reached sharp peaks in mature Ice Ages, just prior to advent of terminations. They proposed that the X-factor is dust deposited on the ice sheets, driving up solar absorption, leading to rapid disintegration of the ice sheets. The dust was generated by desertification of distant marginal regions due to CO_2 starvation and cold, and transported by winds to the ice sheets.

Terminations take place typically in about 5500 years—about half an up-lobe of solar input. An interglacial follows termination. During an interglacial, dust levels are nil and the solar precession curve is on the back half of the up-lobe. After about 5500 years of the interglacial, the solar curve turns downward. Ice begins to slowly accumulate at high latitudes. But sea level remains high until sufficient ice accumulates to reduce sea level. If the duration of an interglacial is measured in terms of when dv/dt turns positive, it might be about 5500 years.

It is interesting to speculate when the next Ice Age might occur in the future. Since it is theorized that deposition of dust produces terminations, the current heavy deposition of soot, ash, dirt and dust on Greenland and other northern sites insures that there will not be another Ice Age in the near future. Increased CO_2 will amplify this conclusion.

South Pasadena, USA

Donald Rapp

Contents

1	Histo	ry and D	Description of Ice Ages	1
	1.1	Discov	ery of Ice Ages	1
	1.2	Descrip	otion of Ice Sheets	7
	1.3	Vegeta	tion During LGM	11
		1.3.1	LGM Climate	11
		1.3.2	Global Flora	14
		1.3.3	Ice Age Forests	20
	1.4	Vegeta	tion and Dust Generation During the LGM	22
		1.4.1	Introduction: Effect of Low CO ₂ on Plants	22
		1.4.2	C3 and C4 Flora Differences	25
		1.4.3	Effects of Low CO ₂ on Tree Lines	25
		1.4.4	Source of the LGM Dust	30
2	Varia	ability of	the Earth's Climate	39
	2.1	Factors	that Influence Global Climate	39
	2.2	Stable	Extremes of the Earth's Climate	43
	2.3	Ice Age	es in the Recent Geological Past	48
3	Ice C	Core Meth	nodology	51
	3.1	History	v of Ice Core Research	51
	3.2	Dating	Ice Core Data	57
		3.2.1	Introduction	57
		3.2.2	Age Markers	58
		3.2.3	Counting Layers Visually	59
		3.2.4	Layers Determined by Measurement	63
		3.2.5	Ice Flow Modeling	65
		3.2.6	Other Dating Methods	68
		3.2.7	Synchronization of Dating of Ice Cores from	
			Greenland and Antarctica	69

		3.2.8	GISP2 Experience	70
		3.2.9	Tuning	70
		3.2.10	Flimsy Logic	71
	3.3	Processi	ng Ice Core Data	73
		3.3.1	Temperature Estimates from Ice Cores	73
		3.3.2	Temperature Estimates from Borehole Models	77
		3.3.3	Climate Variations	80
		3.3.4	Trapped Gases	80
4	Ice Co	re Data		83
7	4 1	Greenla	nd Ice Core Historical Temperatures	84
	4.2	Antarcti	ca Ice Core Historical Temperatures	89
	7.2	4 2 1	Vostok and EPICA Data	89
		422	Homogeneity of Antarctic Ice Cores	89
	43	North-S	outh Synchrony	90
	1.5	431	Direct Comparison of Greenland and Antarctica	20
			Ice Core Records	90
		4.3.2	Sudden Changes	96
		4.3.3	Interpretation of Sudden Change in Terms of Ocean	10
			Circulation	99
		4.3.4	Seasonal Variability of Precipitation	101
	4.4	Data fro	m High-Elevation Ice Cores	101
	4.5	Carbon	Dioxide	102
		4.5.1	Measurements	102
		4.5.2	Explanations	105
	4.6	Dust in	Ice Cores	116
5	Ocean	Sedimen	nt Data	119
5	5 1	Introduc	tion	120
	5.2	Chronol	Ωσ ν	125
	53	Univers	ality of Ocean Sediment Data	129
	5.5	Summar	ry of Ocean Sediment Ice Volume Data	130
	5.5	Compar	ison of Ocean Sediment Data with Polar Ice Core Data	132
	5.6	Historic	al Sea Surface Temperatures	135
	5.7	Ice-Raft	ed Debris	136
	0.1			100
6	Other	Data Sou	urces	139
	6.1	Devil's		139
		0.1.1	Devil's Hole Data	139
		0.1.2	Comparison of Devil s Hole Data with Ocean	1.4.1
		612	Devil's Heley Clobal on Deviced Data?	141
		0.1.3	Comparison of Davilla Hala Data with Vestal Data	143
		0.1.4	Comparison of Devil s Hole Data with Vostok Data	143
		0.1.5	The Continuing Controversy.	145

	6.2	Speleothems in Caves	146	
	6.3	Magnetism in Rocks and Loess	147	
		6.3.1 Magnetism in Loess	147	
		6.3.2 Rock Magnetism in Lake Sediments	148	
	6.4	Pollen Records	148	
	6.5	Physical Indicators	150	
		6.5.1 Ice Sheet Moraines	150	
		6.5.2 Coral Terraces	151	
		6.5.3 Mountain Glaciers	151	
	6.6	Red Sea Sediments	152	
7	Over	view of the Various Models for Ice Ages	155	
	7.1	Introduction	157	
	7.2	Variability of the Sun	158	
	7.3	Astronomical Theory	158	
	7.4	Volcanism	160	
	7.5	Greenhouse Gases	164	
	7.6	Role of the Oceans.	164	
		7.6.1 Glacial-Interglacial Cycles: The Consensus View	164	
		7.6.2 Sudden Climate Change—The Consensus View	168	
		7.6.3 Wunsch's Objections	171	
	7.7	Models Based on Clouds	177	
		7.7.1 Extraterrestrial Dust Accretion	178	
		7.7.2 Clouds Induced by Cosmic Rays	178	
		7.7.3 Ocean–Atmosphere Model	182	
	7.8	Models Based on the Southern Hemisphere	183	
8	Varia	ability of the Earth's Orbit: Astronomical Theory	185	
	8.1	Introduction	185	
	8.2	Variability of the Earth's Orbit	188	
		8.2.1 Variability Within the Orbital Plane	188	
		8.2.2 Variability of the Orbital Plane	192	
	8.3	Calculation of Solar Intensities	192	
	8.4	Importance of Each Orbital Parameter	194	
	8.5	Historical Solar Irradiance at Higher Latitudes		
	8.6	Connection Between Solar Variability and Glaciation/		
		Deglaciation Cycles According to Astronomical Theory	199	
		8.6.1 Models for Ice Volume	201	
		8.6.2 Review of the Imbries' Model	209	
		8.6.3 Memory Model	213	
		8.6.4 Modification of Paillard Model	213	

	8.7	Models Based on Eccentricity or Obliquity8.7.1A Model Based on Eccentricity8.7.2The Middle-Pleistocene Transition (MPT)	220 220 222
9	Compa	arison of Astronomical Theory with Data	227
	9.1	Ice Volume Versus Solar Input	227
	9.2	Spectral Analysis	237
		9.2.1 Introduction	237
		9.2.2 Spectral Analysis of Solar and Paleoclimate Data	241
10	Interg	lacials	249
11	Termi	nations of Ice Ages	257
	11.1	Abstract	262
	11.2	Background	263
	11.3	Terminations	266
	11.4	North or South (or Both)?	271
	11.5	Models Based on CO ₂ and the Southern Hemisphere	276
	11.6	Climate Models for Terminations of Ice Ages	279
	11.7	Model Based on Solar Amplitudes	285
	11.8	Dust as the Driver for Terminations	288
		11.8.1 Introduction	288
		11.8.2 Antarctic Dust Data	289
		11.8.3 Correlation of Ice Core Dust Data with Terminations	290
		11.8.4 Dust Levels on the Ice Sheets	296
		11.8.5 Optical Properties of Surface Deposited Dust	303
		11.8.6 Source of the Dust	304
		11.8.7 Ice Sheet Margins	306
	11.9	Model Based on Solar Thresholds	310
	11.10	The Milankovitch Model Versus the Most Likely Model	313
		11.10.1 Criteria for a Theory	313
		11.10.2 The "Milankovitch" Model	314
		11.10.4 Users and O estimate	310
		11.10.4 Unanswered Questions	318
12	Status	of Our Understanding	319
References 32			327
Ind	ex		343

Abbreviations

AABW	Atlantic Bottom Water
ACP	Age control point
AM	Amplitude modulation
AMO	Atlantic meridional overturning
AMOC	Atlantic meridional overturning circulation
AMSL	Above mean sea level
AWS	Automated weather station
C&L	Chylek and Lohmann
CAS	Central American Seaway
CLIMAP	The "Climate: Long-range Investigation, Mapping, and Prediction" project
CRF	Cosmic ray flux
D-0	Dansgaard–Oeschger events
EAIS	East Antarctic Ice Sheet
ECM	Electro-conductivity measurement
EDC	EPICA Dome C
EDML	EPICA Dronning Maud Land
EEM	Previous interglacial period named after Dutch River
ENSO	El Niño-Southern Oscillation
EOT	Eocene–Oligocene transition
EPICA	European Project for Ice Coring in Antarctica
ERBE	Earth Radiation Budget Experiment
GCM	Global climate model
GCR	Galactic cosmic ray
GICC	Glacial-interglacial CO ₂ cycle
GISP	Greenland Ice Sheet Project
GRACE	Gravity Recovery and Climate Experiment
GRIP	Greenland Ice Core Project
GSLR	Global sea level rise
gya	Billions of years before present
H&A	Hargreaves and Annan

H&W	Huybers and Wunsch (2004)
IPCC	Intergovernmental Panel on Climate Change
IR	Infrared
IRD	Ice rafted debris
kya	Thousands of years before present
kyr	Thousands of years
L&R	Lisiecki and Raymo
L&W	Landwehr and Winograd
L&W	Landwehr and Winograd (2001)
LGM	Last Glacial Maximum
LIA	Little Ice Age
LIG	Last interglacial
LLS	Laser light scattering
M&M	The book by Richard A. Muller and Gordon J. MacDonald: Ice Ages and
	Astronomical Causes, Praxis Publishing (2000)
M&W	McShane and Wyner
MBH	Mann, Bradley and Hughes
MECO	Middle Eocene climatic optimum
MOC	Meridional overturning circulation
MPR	Mid-Pleistocene revolution
MPT	Mid-Pleistocene transition
MWP	Medieval Warm Period
mya	Millions of years before present
NADW	North Atlantic Deep Water
NGRIP	North Greenland Ice Core Project
NH	Northern hemisphere
NHG	Northern hemisphere Glaciation
OLR	Outgoing long-wavelength radiation
PAL	Present atmospheric level
PCA	Principal component analysis
PDB	Crushed belemnite (Belemnitella americana) from the Peedee Formation
	(Cretaceous) in South Carolina
PDO	Pacific decadal oscillation
PETM	Paleocene–Eocene Thermal Maximum
PI	Pre-industrial
RSL	Relative sea level
SH	Southern hemisphere
SMB	Surface mass balance
SMOW	Standard Mean Ocean Water
SST	Sea surface temperature(s)
TIMS	Thermal ionization mass spectrometry

TOA	Top of atmosphere
TSI	Total Solar Irradiance
UWESS	University of Washington Earth and Space Sciences Department
VEI	Volcanic Explosivity Index
W&L	Winograd and Landwehr (1993)
WAIS	West Antarctic Ice Sheet
WB	Wally Broecker
ya	Years ago

List of Figures

Fig. 1.1	Erratic stone	3
Fig. 1.2	Scratched stone	3
Fig. 1.3	Bubble rock, Acadia, Maine	3
Fig. 1.4	Extent of the most recent ice age in North America	6
Fig. 1.5	Glacial striations	7
Fig. 1.6	Extent of the ice sheets 18,000 years ago	8
Fig. 1.7	Extent of the ice sheets 12,000 years ago	9
Fig. 1.8	Extent of the ice sheets 8500 years ago	10
Fig. 1.9	Extent of the ice sheets 7500 years ago	11
Fig. 1.10	Distribution of vegetation in North and Central America at the	
	height of the last Ice Age	15
Fig. 1.11	Distribution of vegetation in North and Central America today	
	if there were no agriculture	16
Fig. 1.12	The Wrangell-Saint Elias ice field on the Alaska-Yukon border	17
Fig. 1.13	Beringia—the connecting link between Siberia	
	and Alaska about 18,000 ya	17
Fig. 2.1	Variation in carbon isotope composition of shallow marine	
	carbonates	47
Fig. 2.2	Global average temperature over the past three million years	49
Fig. 3.1	Isotopic fractionation in vapor and precipitation	53
Fig. 3.2	The sintering process as snow is converted to firn and then	
	on to ice with bubbles of air entrapped	54
Fig. 3.3	Examples of line scan images of ice cores from various depths	61
Fig. 3.4	Close-up examples of line scan images	62
Fig. 3.5	Section of the GISP2 ice core from 1837 to 1838 m deep in	
	which annual layers are clearly visible	63
Fig. 3.6	Section of an ice core drilled in the Kunlun Mountains	
	of Western China.	64
Fig. 3.7	a Layering as evidenced by periodic variations in δ^{18} O	
	(Dansgaard 2005). b Layering of δ^{18} O measurements at Station	
	Crete, Greenland	64

Fig. 3.8	Example of a 1.2-m segment of GRIP data from about 8800 ya	65
Fig. 3.9	Example of data and annual layer markings from visual stratigraphy	
	during the early Holocene	66
Fig. 3.10	Ice particle flow paths	66
Fig. 3.11	Chronology of Antarctic ice core temperatures	
	(Kawamura 2009)	69
Fig. 3.12	Depth-age relationship in GISP2	71
Fig. 3.13	Dansgaard's correlation of δ^{18} O with temperature. The circles are	
	South Greenland and the squares are North Greenland	73
Fig. 3.14	Present day correlations of δ^{18} O or δ D with temperature	75
Fig. 3.15	Relation between isotopic composition of precipitation and	
	temperature in the parts of the world where ice sheets exist	75
Fig. 3.16	Comparison of estimated temperatures at two Greenland sites	76
Fig. 3.17	Filtered isotope data	78
Fig. 3.18	Rate of accumulation increased in the Holocene	79
Fig. 3.19	Difference between isotope ratios in summer and winter	80
Fig. 4.1	Greenland topographical map showing locations of several major ice	
-	core sites. Numbers are elevations in meters	85
Fig. 4.2	GISP2 estimates of global temperatures over the past	
	two centuries. The Medieval Warm Period and Little Ice age are	
	evident.	86
Fig. 4.3	Ice core estimates of global temperatures during the	
	past 12,000 years	86
Fig. 4.4	Greenland temperature history from GRIP over 20,000 years	
	(smoothed data)	87
Fig. 4.5	GISP2 ice core results taken at Greenland summit	
C	over 40,000 years	87
Fig. 4.6	Global temperature estimates from GISP2 ice cores	
	over 100,000 years	88
Fig. 4.7	Greenland temperature history from GRIP (smoothed data) over	
	150,000 years. Interglacial periods are shown by gray shading	88
Fig. 4.8	Antarctica topographical map showing locations of several major	
	ice core sites	90
Fig. 4.9	Vostok ice core data	92
Fig. 4.10	Estimated temperature difference from today at EPICA-Dome	
	C versus age	92
Fig. 4.11	EPICA Dome C temperature data corrected for elevation	
C	changes	93
Fig. 4.12	Comparison of Vostok and Dome C ice core data	93
Fig. 4.13	Comparison of the Vostok and Dome Fuji isotopic	
-	records of δ^{18} O as a function of depth (Watanabe et al. 2003)	94
Fig. 4.14	Comparison of the Vostok and Dome Fuji isotopic	
č	records of δ^{18} O as a function of time	94
Fig. 4.15	Comparison of Greenland and Antarctica isotope records	95

Fig. 4.16	Isotopic and CH ₄ data from Greenland and Antarctica on the GISP2 time scale	96
Fig 4 17	Sudden climate change events at Greenland and Antarctica	98
Fig. 4.18	Vostok (Antarctica) record of CO_2 CH ₄ and temperature	20
11g. 1.10	(from δD)	103
Fig. 4.19	Variation of CO ₂ concentration since the LGM	104
Fig. 4.20	Comparison of ice volume with CO ₂ concentration across the	101
1.8	last termination	105
Fig. 4.21	Variation of δ^{14} C and CO ₂ concentration during	100
8	past 40.000 years.	111
Fig. 4.22	Variation of temperatures and CO ₂ since the LGM	112
Fig. 4.23	Variation of peak solar intensity over the past 40,000 years	113
Fig. 4.24	Sea salt sodium flux measured from the EPICA Dronning	
U	Maud Land ice core and solar forcing	114
Fig. 4.25	Taylor Dome record of atmospheric O_2 over the most recent glacial	
U	termination	115
Fig. 4.26	Phase diagram for CO_2 as a function of temperature	
U U	and pressure.	116
Fig. 4.27	Comparison of dust flux to pCO_2 in ice core	118
Fig. 5.1	Fit of a portion of the δ^{18} O curve to the ice model of Lisiecki	
	and Raymo (2005)	128
Fig. 5.2	Assignment of stages and terminations by H&W. Stages are	
	designated by arrows and terminations are defined by circles	129
Fig. 5.3	Universality of oxygen isotope patterns from forams from around	
	the world	130
Fig. 5.4	SPECMAP showing marine isotope stage numbers	131
Fig. 5.5	δ^{18} O for site 806	131
Fig. 5.6	Isotope data from a stack of 57 records	133
Fig. 5.7	Comparison of variation of δ^{18} O data from two sources	134
Fig. 5.8	Comparison of ocean sediment data with Antarctica EPICA-Dome	
	C data	135
Fig. 6.1	Measured oxygen isotope variability at Devil's Hole	140
Fig. 6.2	Comparison of ages from Devil's Hole with ages	
	from SPECMAP	142
Fig. 6.3	Alignment of Devil's Hole transition points with Vostok transition	
	points.	145
Fig. 6.4	Upper panel: δ^{18} O from GRIP in Greenland. Lower panel: measured	
	susceptibility of lake sediments in France	149
Fig. 6.5	Smoothed data on sea level based on Red Sea sediments and coral	
	terrace data	153
Fig. 7.1	Direct and diffuse solar irradiance measured at Mauna Loa following	
	volcanic eruptions	161
Fig. 7.2	Global temperature change after Toba eruption	163

Fig. 7.3	Near-surface waters flow towards four main deep-water formation regions (yellow ovals)—in the northern North Atlantic,	
	the Ross Sea and the Weddell Sea	166
Fig. 7.4	Schematic of the three modes of ocean circulation that prevailed	
	during different times of the last glacial period	167
Fig. 7.5	Total meridional heat flux of the combined ocean/atmosphere system	
	estimated from Earth Radiation Budget Experiment (ERBE)	
	satellites, direct ocean measurements, and atmospheric contribution	
	as a residual	174
Fig. 7.6	Comparison of Greenland temperature profile with methane	
	profile at Vostok	176
Fig. 7.7	Comparison of radionuclide fluxes with relative amount of ice-rafted	
	debris over the past 12,000 years	181
Fig. 7.8	Comparison of lower troposphere cloud cover anomaly with cosmic	
	ray anomaly over the past two sunspot cycles	182
Fig. 8.1	Motion of Earth about the Sun	189
Fig. 8.2	Variation of obliquity over the past 400,000 years	190
Fig. 8.3	Variation of eccentricity over past 400,000 years	190
Fig. 8.4	Variation of the longitude of perihelion over the	
	past 400,000 years	191
Fig. 8.5	Variability of the tilt of the Earth's orbit plane	193
Fig. 8.6	Variation of daily total solar irradiance with day of the solar year	
	for several northern latitudes	195
Fig. 8.7	Calculated peak summer solar intensity and yearly average of solar	
	intensity at 65°N	197
Fig. 8.8	Calculated peak summer solar intensity at 65°N at three northern	
	latitudes over the past 400,000 years	198
Fig. 8.9	Relative peak summer solar intensity at 65°N	
	over 400,000 years	198
Fig. 8.10	Relative peak summer solar intensity at 65°N	
	over 800,000 years	199
Fig. 8.11	Comparison of peak summer solar intensity to a horizontal surface	
	above the atmosphere at 65°N and 65°S over 400,000 years	199
Fig. 8.12	Weertman's ice sheet model	203
Fig. 8.13	Model for solar input to high latitudes used by Paillard (1998)	210
Fig. 8.14	Dependence of the Imbries' integration on starting values	
	for relative ice volume.	211
Fig. 8.15	Dependence of the Imbries' integration on parameter T	211
Fig. 8.16	Dependence of the Imbries' integration on B	212
Fig. 8.17	Comparison of predicted ice volume from Imbries' theory with	
	measured ice volume	212
Fig. 8.18	Schematic variation of v and dv/dt versus t	215
Fig. 8.19	Schematic representation of (dv/dt) in the transitions	
	to and from Ice Ages.	216

Fig. 8.20	Schematic representation of $v(t)$ in the transitions to and from Ice Ages	217
Fig. 8.21	Comparison of the solar precession curve with sea level over the past	210
Fig. 8.22	Comparison of estimates of ice volume data with mid-summer	219
C	solar input to 65°N	220
Fig. 8.23	Vostok ice core data showing the spacing between rapid	
-	terminations	225
Fig. 9.1	Comparison of Imbries' model with SPECMAP	229
Fig. 9.2	of SPECMAP	231
Fig. 9.3	Comparison of inverse of solar input to 65°N with slope	
C	of HW04 over the past 800,000 years	232
Fig. 9.4	Comparison of inverse solar input to 65°N with slope of HW04 over	
	the past 800,000 years	233
Fig. 9.5	a Comparison of relative ice volume of the SPECMAP with the	
	LR04 stack b Comparison of rate of change of ice volume (dv/dt)	225
F ' 0 (of the LR04 stack with a solar curve.	235
Fig. 9.6	Spectra of sine wave and sawtooth wave with 100 ky periods	240
Fig. 9.7	Three simple functions for spectral analysis	240
Fig. 9.8	Frequency distribution corresponding to functions in Fig. 9.6	241
Fig. 9.9	Function $F(t) = \cos(1.5 t) + \cos(2.5 t)$	241
Fig. 9.10	Spectral distribution of frequencies corresponding to function	2.42
	$F(t) = cos(1.5 t) + cos(2.5 t) \dots$	242
Fig. 9.11	A hypothetical function with quasi-periodic behavior	242
Fig. 9.12	Frequency distribution corresponding to function in Fig. 9.10	243
Fig. 9.13	Spectra for solar intensity at 65°N and 65°S according to M&M	243
Fig 914	Vostok time series (M&M)	244
Fig. 9.15	Spectrum of Vostok deuterium data according to M&M	245
Fig. 9.16	Frequency spectra of EPICA Dome-C Antarctic ice core data	210
1.8	(Petit et al. 1999)	245
Fig 917	Spectrum of SPECMAP according to M&M and Spectrum of	210
118. 9.17	Imbries' ice model.	246
Fig. 10.1	Relative sea level over the past 800,000 years, and a possible range	
-	that defines an interglacial	250
Fig. 10.2	High-resolution carbon dioxide concentration record	252
Fig. 10.3	Relative sea level over the past 800,000 years, and rough estimates	
-	of duration of interglacials, shown by blue dashed lines	254
Fig. 11.1	Temperature deviation in Antarctic ice core and solar input to 65°N	
2	latitude over the past 450,000 years	259
Fig. 11.2	The idealized concept of interruptions in Ice Ages	261

Fig. 11.3	Solar input to 65°N at midsummer shown along with the termination	
F 11.4	ramps.	264
Fig. 11.4	Empirical model for termination ramp	266
Fig. 11.5	Superposition of nine temperature curves around	• • •
-	the last termination	268
Fig. 11.6	Temperature change before and through the current interglacial	269
Fig. 11.7	Temperature change before and through the previous interglacial	270
Fig. 11.8	Variation of CO_2 concentration and dust flux through the last two	
8	terminations.	271
Fig. 11.9	Comparison of dust deposition rate and temperature change through	
0	the last and previous terminations	272
Fig. 11.10	Dust, temperature and solar around the time of the last and previous	272
Fig. 11.11	Palative solar input to 65°N and 65°S showing the two periods	215
11g. 11.11	during which both solar inputs are increasing simultaneously	272
Fig. 11.12	Comparison of modeled rate of dust leading with ice volume	275
Fig. 11.12	Comparison of modeled dust leading with ice volume	201
Fig. 11.15 Eig. 11.14	Comparison of Anteratic ice core date with colculated color input	202
rig. 11.14	at 65°N over 800 000 years	286
Fig 11.15	Antarctic dust concentration in ice cores as measured by the Coulter	200
115. 11.15	method	290
Fig 11.16	Antarctic dust concentration in ice cores as measured by a laser	220
1.8. 11.10	ontical method	291
Fig 11 17	Antarctic temperature (top) Solar intensity at 65°N on	
1.1g. 11.17	June 21 (middle) Dust loading in Antarctica ice core	
	using the Coulter counter (bottom)	292
Fig 11 18	Antarctic temperature (top) Solar intensity at 65°N on	
1.5. 11.10	June 21 (middle) Dust loading in Antarctica ice core using	
	the laser counter (bottom)	293
Fig 11 19	FPICA dust levels at Greenland	296
Fig. 11.19	Ice-margin sampling site in Kronprins Christian Land	307
Fig. 11.20	Photo of ablation zone at Kronprinz Christians Land	308
Fig. 11.21	Pleistocene ice with uniform debris cover and meltwater	500
115. 11.22	streams	309
Fig 11 23	Comparison of solar curves	311
Fig. 11.23	Comparison of smoothed Greenland temperature with yearly solar	511
115. 11.24	input at noon at 65°N over the past 150,000 years	315
Fig. 11.25	Increase in dust precedes rise in CO.	317
1 ig. 11.23	$\frac{1}{1000}$	517

List of Tables

Table 1.1	Reductions in temperature and CO_2 affecting the maximum tree line	
	in the Alps and tropics	26
Table 4.1	Characteristics of Greenland and Antarctica ice sheets	84
Table 4.2	Characteristics of major ice core sites at Greenland	85
Table 4.3	Characteristics of major ice core sites at Antarctica	91
Table 4.4	Dust fluxes in g/m^2 -year estimated by Lambert et al. (2015)	117
Table 7.1	Heating and cooling effects of clouds	177
Table 10.1	Timing of the onset and end of interglacials of the last 800 kyr	
	and their estimated duration according to Tzedakis (2012)	254
Table 10.2	Timing of the onset and end of interglacials of the last 800 kyr	
	and their estimated duration according to Varga (2015)	255
Table 11.1	Durations required for the transition from glacial to interglacial	
	conditions	267
Table 11.2	Relationship of the dust spikes to the ice volume curve in the	
	Ganopolski et al. model	283
Table 11.3	Comparison of Antarctic ice core data with calculated yearly	
	solar input at 65°N over 800,000 years.	287
Table 11.4	Analysis of Fig. 11.17 based on Coulter counter	294
Table 11.5	Analysis of Fig. 11.18 based on laser counter	295
Table 11.6	Accumulation around the time of the LGM	297

History and Description of Ice Ages

Abstract

The existence of past ice ages was discovered by several 19th century geologists from scratch marks on rocks, erratic boulders, moraines, and other physical observations. As early as 1920, Chamberlain provided a map of the North American and Greenland ice sheets at the last glacial maximum that remain quite accurate even today. Two massive ice sheets dominated the northern hemisphere. Nearly a quarter of the earth's surface lay under the weight of a mountain of ice. The Laurentide ice sheet is believed to have reached a height of 12,500 ft. Ice covered nearly 5 million square miles of North America. As the glaciers grew, they drew so much water that the ocean levels dropped more than 100 m. The expansion of the glaciers dramatically affected the distribution and composition of vegetation. Global flora was impacted, by both CO_2 starvation than cold. Deserts expanded and wind-blown dust became prevalent at the last glacial maximum.

1.1 Discovery of Ice Ages

The history of the discovery of the existence of ice ages is summarized nicely in the small book by Woodward (2014). Imbrie and Imbrie (1979) also described this history in their classic book. In addition, Berger (2012) also presented an excellent history.

In the early 1800s, evidence began to emerge of an unusual past. This included the presence of so-called "erratic boulders"—large rounded rocks seemingly placed in inaccessible locations by a giant hand, as well as a multitude of geological evidence of past glacier evidence. However, two factors inhibited the proper interpretation of this as evidence of past ice ages. One was the overhanging influence of the Biblical description of Noah's great flood, which suggested that a great flood might have caused such phenomena. The other was the fact that in geology, there was a debate between advocates of

© Springer Nature Switzerland AG 2019 D. Rapp, *Ice Ages and Interglacials*, https://doi.org/10.1007/978-3-030-10466-5_1



1

slow, gradual evolution of geologic formations versus change via catastrophic events, and at that time the so-called uniformitarian held sway in geology. Other ideas included the rather fantastic notion of transport of large boulders trapped in drift ice.

The Swiss were well positioned to observe the effects of past Ice Ages in the mountains of the Alps. In the 1820s, a Swiss named Venetz showed that the glaciers of the Alps were once far larger than they were at that time. Another Swiss (Charpentier) joined with Venetz in putting forth the proposition that the valleys were once occupied by enormous glaciers, as evidenced by scratch marks on rocks, erratic boulders, moraines, and other observations. A Norwegian (Esmark) found similar evidence in the Fjords. By the early 1830s, these three field investigators had found ample proof of former large-scale glaciation. Nevertheless, the scientific establishment did not give much credence to these findings.

Yet another Swiss, Louis Agassiz, under tutelage of Venetz and Charpentier, became an enthusiast for the glacial theory, and used his high position in Swiss science to promulgate these ideas. Starting in 1840 Agassiz became the main promoter of the glacial theory. In the 1840s he worked with a Scottish geologist (Buckland) to examine the geological evidence in Scotland. As the 1840s began, the glacial theory was still regarded as speculative. It was not until the 1870s that the glacial theory became widely accepted.

In the early 1800s, evidence began to emerge of an unusual past. This included the presence of so-called "erratic boulders"—large rounded rocks seemingly placed in inaccessible locations by a giant hand, as well as a multitude of geological evidence of past glacier evidence. A few geologists of the 19th century noted the presence of large boulders with characteristic scratch marks in the Swiss Alps, as well as scratch marks on the walls of rock in mountains, and suggested that these may have been generated by huge ancient glaciers that covered the mountains. The three main sources of evidence were: (1) grooves and scratches on rocks in place, and on boulders shoved along under the ice, (2) extensive unstratified deposits known as "till" traceable to glacier action, and (3) transported material (boulders) that could only have been delivered by ice (not water).

Prior to the implementation of ice core drilling and use of ocean and lake sediments to infer historical temperatures tens or hundreds of thousands of years ago, geologists had to rely on their observations of rocks and strata for guidance. Three books written around the end of the 19th century provide good insights into what was known prior to modern techniques for estimating historical temperatures. One of these books, Geike (1894) provided Figs. 1.1 and 1.2.

"Bubble Rock" in Maine is a favorite subject for photographers (see Fig. 1.3).

According to another early book, Wright (1920) showed that rocks with scratches and striations longitudinally along their longest diameters are evidence of glacial action:

It is easy to see that the stones of all sizes, while being dragged along underneath the ice, would be held in a comparatively firm grasp as to be polished and striated and scratched in a peculiar manner. On the shores of bays and lakes and in bottoms of streams we find that the stones are polished and rounded in a symmetrical manner, but are never scratched. The mobility of water is such that the edges and corners of the stones are rubbed together by

Fig. 1.1 Erratic stone (Geike 1894)







Fig. 1.3 Bubble rock, Acadia, Maine (http://flickr.com/ photos/iamtonyang/29259194/)



forces acting successively in every possible direction. But in and under the ice the firm grasp of the stiff semi-fluid causes the stony fragments to move in a nearly uniform direction, so that they grate over the underlying rocks like a rasp From the stability of the motion of such a

substance as ice there would ... result grooves and striation both on the rocks beneath and on the boulders and pebbles that, like iron plowshares, are forced over them. Scratched surfaces of rock and scratched stones are therefore, in ordinary cases, most trustworthy indications of glacial action. The direction of the scratches upon these glaciated boulders and pebbles is, also worthy of notice. The scratches upon the loose pebbles are mainly in the direction of their longest diameter—a result that follows from a mechanical principle, that bodies forced to move through a resisting medium most swing around so as to proceed in the line of least resistance. Hence the longest diameter of such moving bodies will tend to come in line with the direction of the motion.

However, Wright (1920) cautioned:

A scratched surface is, however, not an infallible proof of the former presence of a glacier where such a surface is found, or, indeed, of glacial action at all. A stone scratched by glacial forces may float away upon an iceberg and be deposited at a great distance from its home. Indeed, icebergs and shore-ice may produce, in limited degree, the phenomena of striation that we have just described.

Wright (1920) went on to say that although longitudinal striations can be caused by factors other than moving ice, these can by identified by the informed observer:

Stones are also striated by other agencies than moving ice. Extensive avalanches and landslides furnish conditions analogous to those of a glacier, and might in limited and favorable localities simulate its results. In those larger geological movements, also, where the crust of the earth is broken and the edges of successive strata are shoved over each other, a species of striation is produced. Occasionally this deceives the inexperienced or incautious observer. But by due pains all these resemblances may be detected and eliminated from the problem, leaving a sufficient number of unquestionable phenomena due to true glacial action.

Wright (1920) also made the point that deposits left by moving water are always stratified:

A second indubitable mark of glacial motion is found in the character of the deposit left after the retreat of the ice. Ice and water differ so much from each other in the extent of their fluidity, that there is ordinarily little danger of confusing the deposits made by them. A simple water deposit is inevitably stratified. The coarse and fine material cannot be deposited simultaneously in the same place by water alone. Along the shores of large bodies of water the deposits of solid material are arranged in successive parallel lines, the material growing finer and finer as the lines recede from the shore. The force of the waves is such in shallow water that they move pebbles of considerable size. Indeed, where the waves strike against the shore itself, vast masses of rock are often moved by the surf. But, as deeper water is reached, the force of the waves becomes less and less at the bottom, and so the transported material is correspondingly fine, until, at the depth of about seventy feet, the force of the waves is entirely lost; and beyond that line nothing will be deposited but fine mud, the particles of which are for a long while held in suspension before they settle.

In the deltas of rivers, also, the sifting power of water may be observed. Where a mountain-stream first debouches upon a plain, the force of its current is such as to move large pebbles, or boulders even, two or three feet in diameter. But, as the current is checked, the particles moved by it become smaller and smaller until in the head of the bay, or in the broad current of the river which it enters, only the finest sediment is transported. The difference

between the size of material transported by the same stream when in flood find when at low water is very great, and is the main agent in producing the familiar phenomena of stratification. During the time of a flood, vast bodies of pebbles, gravel, and sand are pushed out by the torrent over the head of the bay or delta into which it pours; while during the lower stages of water only fine material is transported to the same distance; and this is deposited as a thin film over the previous coarse deposit. Upon the repetition of the flood another layer of coarser material is spread over the surface; And so, in successive stages, is built up in all the deltas of our great rivers a series of stratified deposits. In ordinary circumstances, it is impossible that coarse and fine material should he intermingled in a water deposit without stratification. Water moving with various degrees of velocity is the most perfect sieve imaginable; so that a water deposit is of necessity stratified.

By contrast, deposits left by moving ice are not stratified:

It is evident that ice is so nearly solid that the earthy material deposited by it must be unassorted. The mud, sand, gravel, pebbles, and boulders, dragged along underneath a moving stream of ice, must be left in an unstratified condition—the coarse and the fine being indiscriminately mingled together. This is the character of the extensive deposits of loose material that cover what we designate as a glaciated region [In such an] unstratified deposit, a variety of materials is mingled that were derived from rocks both of the locality and from far-distant regions. Moreover, the pebbles in this deposit are the most of them polished and scratched after the manner of those which we know to have been subjected to glacial action.

Finally, Wright (1920) discussed the fact that the southern margin of the region where unstratified deposits containing striated stones and transported material was exceedingly irregular in two respects. The southern edge of these deposits does not follow a straight east-and-west line, but in places withdraws to the north (crenate character), and in other places extends lobe-shaped projections far to the south (serrate character). According to Wright, it was the crenate character of its southern border that was of most significance. Wright emphasized that the southern border, with its indentation, and projections was not determined by any natural barrier based on the geography of the region, but instead was determined by "the irregular losses in momentum such as would take place in a semi-fluid moving in the line of least resistance from various central points of accumulation."

In the late 19th century, Thomas C. Chamberlain (as reported by Geike 1894) reviewed the geological evidence for glacial phenomena on the Earth's surface, that existed prior to acquisition of ice core and benthic data on past Ice Ages. In North America, it was found that a tract of about 4,000,000 square miles had been overspread by glaciers, and nearly one-half of North America was covered with drift deposits. He mentioned concerns of the doubters but concluded: "the uncompromising evidence of the deposits themselves and by the ice-grooved rock floor on which these rest, seems to compel acceptance of the glacial theory." Chamberlain concluded that the extent of the ice sheet was roughly as shown in Fig. 1.4. Note the three epicenters for ice sheet formation.

These descriptions represent only a fraction of the ample evidence available to late 19th century geologists that there was a previous Ice Age, although the possible existence of multiple historical Ice Ages could only be conjectured.



Fig. 1.4 Extent of the most recent ice age in North America (Geike 1894)

The University of Washington Earth and Space Sciences Department (UWESS) produced a number of excellent presentations on Ice Ages that are very descriptive and instructive. The entire structure of the great valley in Yosemite National Park is presented as an example of a classic alpine glaciated landscape.

Glacial erosion occurs by abrasion, crushing and fracturing, and quarrying of joint blocks. Ice is not hard enough to abrade rocks, but rock fragments imbedded in the base of the glacier can abrade rocky terrain below, leaving characteristic striations (see Fig. 1.5).

The UWESS described how glacier action can pluck large blocks leaving characteristic scalloped terrain. In addition, the UWESS provided many more examples and illustrations of past glacial action.



Fig. 1.5 Glacial striations (UWESS)

1.2 Description of Ice Sheets

Carroll et al. (2001) reported:

... during the late glacial period, the Wisconsin, two large ice caps, the Laurentide glacier in the East and Cordellian Glacier in the West, dominated northern North America. Nearly all of Canada lay under the two massive glaciers, which extended into the northern regions of the United States and into the southern one-third of Alaska.

These two massive ice sheets were part of an even larger system of ice that dominated the northern hemisphere. Nearly a quarter of the earth's surface lay under the weight of a mountain of ice. The Laurentide ice sheet is believed to have reached a height of 12,500 ft (Hughes 1987). Ice covered nearly 5 million square miles of North America. As the glaciers grew, they drew more than 50% of the Earth's available water, affecting precipitation The ocean levels dropped, exposing what we call the Continental Shelfs. The expansion of the glaciers dramatically affected the distribution and composition of vegetation.

The leading edge of the glacier in the United States is believed to have been over a mile high (Hughes 1987). Nothing could stand in the way of this massive ice field as it pushed south, grinding over mountains and depressing the land under its massive weight. Over the ice caps, a huge high-pressure system pushed the polar jet stream southward, dominating weather patterns over much of the northern hemisphere The ice sheets influenced temperatures far to the south, and both vegetation and wildlife retreated in its front.



Fig. 1.6 Extent of the ice sheets 18,000 ya (Carroll et al. 2001)

Carroll et al. (2001) provided Figs. 1.6, 1.7, 1.8 and 1.9. Figure 1.6 shows the extent of the North American ice sheets at the Last Glacial Maximum (LGM). Note that the ice sheets covered all of the land above 50°N except for Alaska, and penetrated down to 40°N in the Northeastern United States. Alaska was cold enough to support ice sheets but did not receive enough precipitation. Figures 1.7, 1.8 and 1.9 show the successive depletion of the ice sheets during termination.

Hughes et al. (2015) presented a new time-slice reconstruction of the Eurasian ice sheets documenting the spatial evolution of these interconnected ice sheets every 1000 years from 25 to 10 thousand years ago (kya), and at four selected time periods back to 40 kya. At the height of the last glacial period ice sheets extended over (1) all of Greenland including the margins, out to small areas of the surrounding ocean, (2) the Barents Sea extending from Svalbard to the Kara Sea, (3) over all of Ireland and Scotland and the surrounding seas, and (4) over all of Norway, Sweden, Finland, and down to Lithuania.