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Indrani Roy

Climate Variability and Sunspot Activity

Analysis of the Solar Influence on Climate



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Analysis of the Solar Influence on Climate



Indrani Roy Mathematics and Physical Sciences University of Exeter, College of Engineering Exeter, UK

ISSN 2194-5217 ISSN 2194-5225 (electronic) Springer Atmospheric Sciences ISBN 978-3-319-77106-9 ISBN 978-3-319-77107-6 (eBook) https://doi.org/10.1007/978-3-319-77107-6

Library of Congress Control Number: 2018934396

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Dedicated to my late parents Sri Barin Kar & Smt Jyotsna Kar

Contents

Part I Climatology, General Circulation, Climate Variability and Stratosphere-Troposphere Coupling

1	Clin	natology	and General Circulation	3
	1.1		cology: SLP and SST	3
	1.2		al Circulation	(
		1.2.1	Meridional Circulation	(
		1.2.2	Jet Formation: Thermal Wind Balance Relationship	8
		1.2.3	Walker Circulation	9
2	Maj	or Mode	es of Variability	1
	2.1		ility in the Troposphere	12
		2.1.1	El Niño and Southern Oscillation (ENSO)	12
		2.1.2	North Atlantic Oscillation (NAO)	14
		2.1.3	Arctic Oscillation (AO) and Antarctic	
			Oscillation (AAO)	1:
		2.1.4	Pacific Decadal Oscillation (PDO)	1
		2.1.5	Atlantic Multidecadal Oscillation (AMO)	1
		2.1.6	Indian Summer Monsoon (ISM)	1
		2.1.7	Indian Ocean Dipole (IOD)	2
	2.2	Variab	ility in the Stratosphere	2
		2.2.1	Quasi-Biennial Oscillation (QBO)	2
		2.2.2	Stratospheric Sudden Warming (SSW)	2
	Refe	renes		2
3	Stratosphere-Troposphere Coupling			
	3.1			29
	3.2	Discussion with Schematic		3
	3.3	Strength of Westerly: Solar Influence		
	3.4	Role of Zero Wind Line: QBO Influence		
	3.5	Sun, QBO and Polar Temperature in North Pole		
	3.6	Compo	osites of Time Height Development of NAM	34

viii Contents

	3.7	Annular Modes Pattern Similar	3.			
	3.8	Solar Influence: Polar Vortex and Tropical Lower Stratosphere	3.			
	3.9	Solar Influence: Tropical Lower Stratosphere to Troposphere	3			
	Refe	erences	3			
4	Tele	connection Among Various Modes	3			
	4.1	Polar Vortex, QBO and ENSO	3			
	4.2	Polar Vortex and ENSO	4			
	4.3	ENSO and Polar Troposphere	4			
	4.4	ENSO, Polar Annular Modes and JET	4			
	4.5	ENSO Teleconnections	4			
		4.5.1 El Niño (Warm) and La Niña (Cold) Definition	4			
		4.5.2 El Niño or La Niña?	4			
		4.5.3 ENSO Seasonal Locking	4			
		4.5.4 Potential Problems with SST Data	4			
		4.5.5 Indian Summer Monsoon and Walker Circulation	4			
		4.5.6 Different Types of ENSO	4			
		4.5.7 Homogeneous Monsoon Region	4			
		4.5.8 ENSO ISM Correlation	4			
		4.5.9 SST Composites: EN vs. LN	4			
		4.5.10 ISM ENSO Teleconnection Compositing: EN vs. LN	5			
		4.5.11 Rainfall in South America ENSO				
		(Different Types) Teleconnection	5			
		4.5.12 Summary: ENSO and Teleconnections	5			
	Refe	erences	5			
5		Solar Influence Around Various Places: Robust Solar				
		al on Climate	5			
	5.1	Signal on Sea Level Pressure (DJF) Using Multiple				
		Linear Regression	5			
		5.1.1 Method of Multiple Regression Analysis	5			
	5.2	Solar Signal Around Aleutian Low (AL) and Pacific				
		High (PH)	5			
	5.3	Solar Influence: Tropical Pacific SST	5			
	5.4	ENSO and Sun Phase Locking	5			
	5.5	Solar Signal in Tropical Pacific SST Using Compositing	5			
		5.5.1 Method of Solar Peak Year Compositing	5			
	5.6	Observation: Annual Mean Temperature	5			
	Refe	erences	6			
6		l Solar Irradiance (TSI): Measurements				
		Reconstructions	6			
	Pefe	prances	6			

Contents ix

Part	t II A	Atmosph	ere-Ocean Coupling and Solar Variability			
7	Ocea	n Coupli	ing			
	7.1	Shallow	Overturning Circulation			
	7.2	ENSO.				
		7.2.1	ENSO, Thermocline and Upper Ocean Heat Content			
		7.2.2	ENSO and Delayed Oscillator Theory			
		7.2.3	ENSO and Shallow MOC in Tropical Pacific			
		7.2.4	Pycnocline Convergence vs. SST			
		7.2.5	Abrupt Rise in Temperature During 1977–1998			
	Refe	rences				
8	The Sun and ENSO Connection-Contradictions					
	and	and Reconciliations				
	8.1		ignal and ENSO			
	8.2		liction (I): Solar Signal on Tropical Pacific			
			tive Solar Years and ENSO			
	8.3		liction (II): Solar Signal on Tropical Pacific			
			Niño or La Niña			
	8.4	Propose	ed Mechanism: Earlier Period			
	8.5	Propose	ed Mechanism: Later Period			
	8.6	Contrac	lictions and Reconciliations			
	Refe	rences				
	A De	A Debate: The Sun and the QBO				
	9.1	Data Ar	nalysis: Solar and QBO Separately			
	9.2	Polar Temperature During JF with Respect to QBO				
		(40 hPa) and F10.7				
	9.3	Polar Temperature During JF for QBO (30 hPa) and F10.7				
	9.4	Time Series of QBO at Different Height and EOF Analysis				
	9.5		ned Effects: Solar with QBO			
	9.6	Summary				
	Refe	ferences				
0	Sola		ce: 'Top Down' vs. 'Bottom Up'			
	10.1		ifluence: 'Top Down'			
		10.1.1	Solar Influence: 'Top-Down' – via Polar Vortex and Lower Stratosphere			
		10.1.2	Solar Influence: 'Top-Down' – via Lower			
		- · · · -	Stratosphere to Troposphere			
		10.1.3	Solar Influence: 'Top-Down' – via Stratospheric			
			Polar Vortex to Polar Troposphere			
	10.2	Solar In	ifluence: 'Bottom-Up'			
			muchec. Bottom op			

x Contents

11	An O	verview	of Solar Influence on Climate	117		
	11.1	Introduc	ction	117		
		11.1.1	Methodology	119		
	11.2					
	11.3	Results	Results Text			
		11.3.1	Atmosphere Only: Sun and QBO	133		
		11.3.2	Ocean (Only Pacific) and Atmosphere Coupling:			
			Sun, QBO and ENSO	134		
		11.3.3	Atmosphere and Ocean (Only Pacific) Coupling:			
			Sun, QBO, ENSO and Climate Change	134		
	11.4	Discuss	ion	134		
	References					
Part	Ш	Other M	ajor Influences on Climate			
12	Sun:	Atmospl	here-Ocean Coupling – Possible Limitations	143		
	12.1	Sun: At	mosphere-Ocean Coupling 'Top-Down'			
		vs. 'Bot	ttom-Up' Mechanism: a Case Study	143		
	12.2	Sun: At	mosphere-Ocean Coupling – Limitations			
		of Peak	Year Compositing	144		
		12.2.1	Solar Cycle Signals in Peak Year Compositing			
			for SLP: a Case Study	144		
		12.2.2	Solar Cycle Signals in Peak Year Compositing			
			for Indian Summer Monsoon: a Case Study	148		
	12.3	B Difference in Winter Surface Climate Between Solar				
		Minimum and Maximum				
	12.4					
	MLR Technique			151		
		12.4.1	Sun (Using SSN) and NAO in Two Different			
			Time Periods (1856–1977) and (1878–1997)	151		
		12.4.2	Sun (SSN) and NAO Longer Period (1870–2010)	152		
		12.4.3	Sun (SSN) and NAO Lag Relationship	153		
	12.5		nd PDO Relationship	154		
	References 1:					
13	The A	The Arctic and Antarctic Sea Ice				
	13.1		Sea Ice: Last Few Years	158		
	13.2	8				
	13.3					
	13.4					
	Refer	ence		163		
14	CMIP5 Project and Some Results					
	14.1		Climate Models (GCMs): Basic Equations	165		
	14.2	CMIP5	Project	166		

Contents xi

	14.3	Experir	nents: Historical and RCP (Representative		
		Concen	tration Pathway) Scenarios	167	
	14.4	Some C	CMIP5 Models	168	
	14.5	Temper	rature in CMIP5 and Observation	168	
	14.6	Indian S	Summer Monsoon (ISM) and ENSO in CMIP5 Models	170	
		14.6.1	CMIP5 Models for ISM Are Performing Well	170	
		14.6.2	CMIP5 Models for ISM Not Performing Well	171	
		14.6.3	Models: CMIP5, AMIP5 and High Top, Low Top	177	
		14.6.4	Precipitation Composites- El Niño:		
			(CMIP5 vs. AMIP5)	179	
		14.6.5	Changes in ENSO Variability 2050–2100		
			in CMIP3 Experiments	181	
		14.6.6	Stratospheric Features in CMIP5: Low and High		
			Top Models	181	
		14.6.7	Simulated and Observed Stratospheric Temperature	181	
	Refer	ences		184	
15	Gree	een House Gas Warming			
15	15.1		f Radiation	187 187	
	15.2		adiation vs. Terrestrial Radiation	187	
	15.3		on Transmitted by the Atmosphere	107	
			nospheric Windows	188	
	15.4		tion: Water Vapour and CO ₂	188	
	15.5	_	a Greenhouse Gas.	190	
	15.6		rature and CO ₂ : 400,000 Years	190	
	15.7		Temperature Change in the Last 2000 Years	190	
	15.8		ve Forcing	191	
	15.9		Energy Balance	194	
	Refer			195	
16	Volca	nic Infl	uences	197	
	16.1		Cooling Effect	197	
	16.2		ces of Volcanic Eruption	198	
	16.3		of Large Eruptions on Weather and Climate	199	
	16.4		Varming Associated with Large Eruptions	200	
	16.5		olcano and ENSO	201	
	Refer			202	
17	Ozon	e Deplet	tion in the Stratosphere	203	
	17.1		Hole and Montreal Protocol	203	
	17.2		Hole Animation	204	
	17.3		ouse Gases and Ozone in Model	204	
	Refer	ence		205	

xii Contents

18	Influence of Various Other Solar Outputs		
	18.1	Mechanisms	208
	18.2	Other Influences, e.g. Galactic Cosmic Rays	209
	18.3	Sunspot vs. Galactic Cosmic Ray (GCR)	209
	Refe	rences	211
Few	Quest	ions and Exercises for Students	213
Fur	ther R	eading	215
Ind	e x		217

Sources to Figures

- Fig. 1.1 Background maps generated via http://cci-reanalyzer.org/, using NCEP/NCAR Reanalyses v1 data, ClimateReanalyzer.org, University of Maine, Climate Change Institute
- Fig. 1.2 Plots generated using the data from NOAA/OAR/ESRL PSD, Boulder, Colorado, USA, from their website at (http://www.esrl.noaa.gov/psd/)
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xiv Sources to Figures

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Sources to Figures xv

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xvi Sources to Figures

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Introduction

The energy to drive the global atmospheric circulation originates from the Sun. During higher solar activity, the Earth is subjected to enhanced solar irradiance. However, the connection between the Sun and climate has been a puzzle. This is because during a typical 11-year cycle the solar energy output varies by only about 0.1% (Lean and Rind 2001). Based on energy considerations, such a change is too small to produce significant changes in surface conditions. The degree to which variation in solar output affects climate has been the topic of extensive research for a long time.

But nowadays, there is a common consensus that the variations in the UV part of the spectrum between solar maxima and minima (6–8%) leads to more ozone and warming during solar maxima in the upper stratosphere (Crooks and Gray 2005; Hood 2004 and Haigh 1994). Such modulations can impact tropospheric climate. The primary motivations for exploring the areas of solar 11-year cycle variability on climate are the following:

- The Sun is the principal source of energy in the earth; it causes day/night and seasons. There are reasons to believe that solar variability can influence the climate.
- As the Sun follows the 11-year cyclic variability, the knowledge about the Sunclimate relationship on that scale can be used for future climate prediction purpose.
- However, regarding energy output, there is only 0.1% change from solar minimum to maximum years of the 11-year variability, which is too negligible to influence the climate of the earth.
- Moreover, a significant signal is detected on some meteorological parameters, but most of the time that is regionally different. It can also vary with time periods. Unless there are mechanisms to support such behaviour, it may be a mere coincidence.

Other solar-related drivers are also found to influence the climate of earth, for example, geomagnetic activity, total solar irradiance (TSI) and solar cosmic or magnetospheric energetic particle precipitation, and are discussed briefly in the final chapter

xviii Introduction

of the book. However, the current analysis mainly focuses on solar 11-year cyclic variability (as detected for Sunspot number (SSN)). It is because the knowledge about the Sun–climate relationship on that scale can be used for future projection purposes and hence has implications for improved climate prediction. Moreover, unlike other solar-related drivers, there is a very well-accepted mechanism for the SSN, based on solar UV-related variability.

This work is the collection of lecture notes as well as synthesised analyses of published papers on the described subjects. It is divided into three parts: Part I discusses general circulation, climate variability, stratosphere–troposphere coupling and various teleconnections. Part II mainly explores the area of different solar influences on climate and also discusses about ocean –atmosphere coupling. But without a prior knowledge of other important influences on the earth's climate, the understanding of the actual role of the Sun remains incomplete. Hence, Part III covers burning issues such as greenhouse gas warming, volcanic influences, ozone depletion in the stratosphere and Arctic and Antarctic sea ice. At the end of the book, there are few questions and exercises for students.

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Part I Climatology, General Circulation, Climate Variability and Stratosphere-Troposphere Coupling

Abstract Part I initially focused on basic definitions of Climatology, General Circulation, Climate Variability and Stratosphere Troposphere Coupling. There was a discussion on Climatology of Sea Level Pressure (SLP) and Sea Surface Temperature (SST) which was followed by defining Hadley and Walker circulation. In the subsequent chapter, major modes of climate variability are included with their spatial characteristic and temporal behavior. It is then followed by an overview of Stratosphere-Troposphere coupling. Later on, there is a discussion based on teleconnection among various modes. Discussion on how robust solar influences around different places are detected is included afterwards. Finally, there is a discussion on Total Solar Irradiance (TSI) and how it is reconstructed.

Keywords Climatology · Modes of variability · Stratosphere-Troposphere Coupling · Total Solar Irradiance (TSI) · General circulation

Chapter 1 Climatology and General Circulation



Abstract This chapter focused on basic definitions of climatology and general circulation. There was a discussion on climatology of sea level pressure (SLP) and sea surface temperature (SST) which was followed by defining north-south Hadley and east-west Walker circulation. It also defined and described Ferrel cell, Polar cell and various jets. It explained thermal-wind balance relationship and its relevance to jet formation.

Keywords Climatology · Sea Level Pressure (SLP) · Sea Surface Temperature (SST) · Hadley Cell · Ferrel Cell · Polar Cell · Walker Circulation · Subtropical Jet · Thermal Wind Balance · Aleutian Low · Icelandic Low · Azore High · Intertropical Convergence Zone (ITCZ)

To begin with, I start with the definition of weather and climate. Weather is the changing atmospheric conditions (as they affect people), such as precipitation, mist, fog, etc. or meteorological elements like temperature, humidity, winds, etc. On the other hand, the climate is weather averaged over an extended period (say, 30 years).

Below is a brief description about climatology (30 years average) of sea surface temperature (SST) and sea level pressure (SLP) followed by a description of the general circulation.

1.1 Climatology: SLP and SST

The climatology of SLP (Fig. 1.1) and SST (Fig. 1.2) is discussed during two different seasons, the boreal winter and summer. Boreal winter or summer means Northern Hemispheric (NH) winter or summer season. In contrast, austral winter or summer means Southern Hemispheric (SH) winter or summer.

The mean SLP during boreal winter (December-January-February, DJF) differed to that from boreal summer (June-July-August, JJA) and illustrated in Fig. 1.1a, b, respectively. The seasonal variations in SLP are most apparent in the NH. During winter the high-latitude oceans are characterised by low-pressure centres with the Aleutian Low (AL) and Icelandic Low centre in the northern margins of the Pacific

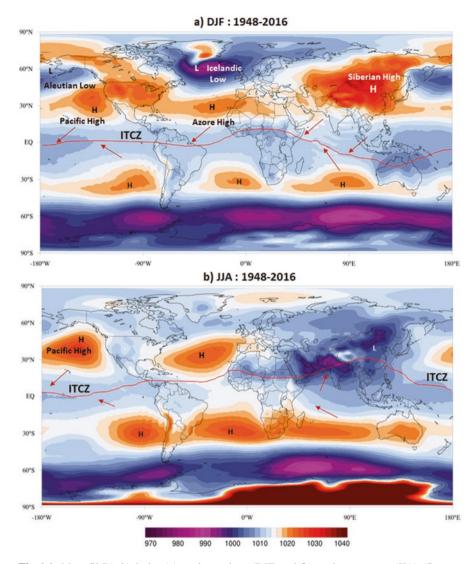


Fig. 1.1 Mean SLP(mb) during (a) northern winter (DJF) and (b) northern summer (JJA). (Source: Background maps generated via http://cci-reanalyzer.org/, using NCEP/NCAR Reanalyses v1 data, ClimateReanalyzer.org, University of Maine, Climate Change Institute)

and Atlantic Oceans, respectively, whereas a high-pressure centre lies over Asia. During summer, the land-sea pressure contrast is reversed in midlatitudes, with the highest pressures over the oceans and the lowest pressures over the land areas. It is seen from that figure that most high-pressure regions persist throughout the year. Although, the Pacific High (PH) and Azore High are weaker in the winter than summer. The dominant low-pressure feature during NH summer is centred over Asia at about 30°N and associated with the Asian summer monsoon. Movement of the intertropical convergence zone (ITCZ), further north around the Indian Ocean

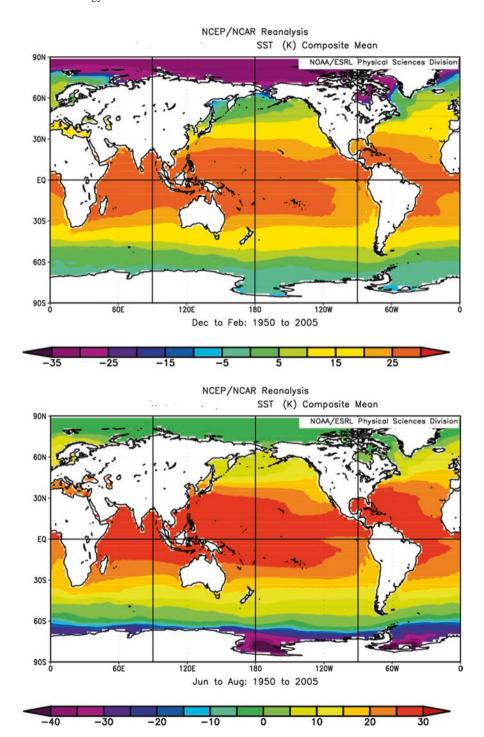


Fig. 1.2 NOAA extended SST (°C) composite mean. (Source: Plots generated using the data from NOAA/OAR/ESRL PSD, Boulder, Colorado, USA, from their website at (http://www.esrl.noaa.gov/psd/))

region during summer, (shown in Fig. 1.1b compared to 1.1a) is clearly noticeable – which is responsible for monsoon around the Indian subcontinent region, causing heavy rainfall.

During winter (shown in Fig 1.1a), the AL and Icelandic Low are well developed. The AL extends from the Aleutian Islands into the Gulf of Alaska, and much stormy weather and precipitation in the Western USA are associated with its movement, whereas the strong circulation around the Icelandic Low produces northerly winds and cold weather in the eastern section of North America. Like the AL, ITCZ and PH are frequently in use in our subsequent discussion; we mention their mean position. The AL covers ~120°W-130°E, 35°N-70°N, whereas the PH centres between 20°N and 50°N, 100°W-140°E.

Mean SSTs during boreal winter (represented here by December–January–February (DJF)) and summer (represented here by June–July–August (JJA)) are illustrated in Fig. 1.2a, b, respectively. During both the seasons, the equator is warmer than the pole. Such temperature gradient is responsible for driving the meridional heat transport through different circulation cells and is described below in the chapter on general circulation.

1.2 General Circulation

Earth's equatorial regions receive more heat than the areas nearer to the pole, which experiences a net deficit (also shown in Fig. 1.2). To prevent the equatorial region from getting warmer and the polar region getting cooler, there must be a transport of heat from the equatorial region towards the pole. Such transport of heat is associated with following meridional circulation cells (shown in Fig. 1.3).

1.2.1 Meridional Circulation

The north south meridional circulation comprises of three cells; Hadley cell, Ferrel Cell and Polar cell.

Hadley Cell Persistent surface heating around the equator results in a rising and poleward moving air at the equator; as this air moves poleward, it cools radiatively and sinks near 30° latitude; it finally returns towards the equator, at low levels. This meridional circulation cell is called the Hadley cell. It is a thermally direct circulation where heat from the sun is converted to motion that transports energy from warm to cold regions. The high-pressure bands around 30° latitude and the low pressure near the equator are signatures of the sinking and rising portion of this Hadley cell, respectively (shown in Fig. 1.3).

Ferrel Cell In midlatitudes the cell that circulates in the opposite direction to the Hadley cell is known as the Ferrel cell, and the overall movement of surface air in